

STUDY OF LATE EOCENE BIVALVES FROM BUDA HILLS

by

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Abstract

This paper completes our knowledge of Late Eocene bivalves of Buda Hills. The 370 specimens, collected from 5 localities, belong to 21 species of 15 genera. This study was carried out from palaeoecological, palaeobiogeographical and biostratigraphical points of view. The five rock types containing this bivalve assemblage, correspond to four different palaeoecological environments. Ascending in the sequence the bivalve assemblage of different rock types indicates increase of sea depth. Results of palaeobiogeographical studies are in good agreement with survey supposing South Alpine connections. This fauna is typical of Late Eocene age and species confined only to the Middle Eocene or the Oligocene were not recorded.

Introduction

Although the Late Eocene mollusc fauna of Buda Hills has been thoroughly investigated for some 120 years, no comprehensive monograph was published. It was high time to carry out an up-to-date study of this fauna from palaeoecological, palaeobiogeographical and biostratigraphical points of view.

KOCH (1872) gave a description of Várerdő Hill sequence at Solymár and also mentioned its fauna.

HOFMANN's activity (1873) was very important in contributing to the knowledge of bivalves of Buda Hills. He gave the first descriptions of these bivalves while he studied the Paleogene deposits of Buda region.

LŐRENTHEY (1897) dealt mainly with Tertiary crustaceans, but he also published a large list of molluscs.

LŐVY (1928) not only listed the fauna of Martinovics Hill but gave ecological evaluations as well.

BOKOR (1939) investigated the deposits of the western margin of Buda Hills and mentioned eleven bivalve species belonging to Eocene.

There was much less contribution to the knowledge of Late Eocene bivalves of this region later in this century.

In the last few decades DUDICH (1959), MONOSTORI (1965, 1967), WEIN (1977) and KÁZMÉR (1982) investigated the geology, mineralogy, microfacies and tectonics of Upper Eocene formations. BÁLDI (1986) published new data on the sedimentology, the palaeontology and the age of Buda Marl.

Specimens, studied in this paper, are from the Museum of the Hungarian Geological Institute, the Department of Palaeontology of Eötvös University and from supplementary collections by the author.

Geological setting

According to BALÁZS et al. (1981) the Upper Eocene of Buda Hills belongs to the epicontinental terrigenous-carbonate development. These sediments discordantly overlie on Triassic dolomites, limestones and Middle Eocene Miliolina limestone.

Late Eocene sea transgressed into Buda region in two steps (DUDICH, 1959). The sequence starts with conglomerate and the *Nummulites* - coralline algae - *Discocyclus* limestone gradually took place. Bryozoan Marl covers this limestone (KÁZMÉR, 1982) either with successive transition on the Mátyás Hill or discordantly with basal conglomerates on the Martinovics Hill. The next unit is Buda Marl, which underlies the following Oligocene sequences or the Neogene and the Quaternary.

The limestone containing *Nummulites fabianii* indicates Priabonian age. According to KOPEK et al. (1969) it is the *N. fabianii* - *Discocyclus* level (no. XIV) of Transdanubian Central Range.

BÁLDI-BEKE (1970) put Bryozoan and Buda Marls into the *Insthomolitus recurvus* zone of Priabonian on the basis of similar nannoplankton assemblages.

Localities

Studied specimens were derived from five localities (Fig. 1):

1. Várerdő Hill at Solymár (Fig. 2).
2. Szépvölgy
3. Martinovics Hill
4. Vár Hill
5. Németsvölgyi road.

Specimens were collected layer by layer except for the eastern quarry of Mátyás Hill, where the collection was from the scree of the silicified marl. Unfortunately the bulk of the specimens were gathered decades ago without exact identification of the place or the layer. Due to the same reason individuals from Szépvölgy were considered as one unit.

Sixty-three specimens from Várerdő Hill, 189 from Szépvölgy, 90 from Martinovics Hill, 12 from Vár Hill and 16 from Németsvölgyi road were examined.

1. Várerdő Hill at Solymár

Eocene and Oligocene deposits crop out in the valley of Jegénye Creek. The detailed geological description is given by MONOSTORI (1967). *Nummulites* - *Discocyclus* limestone underlies the Oligocene Hárshegy Sandstone by angular unconformity (Fig. 3).

According to KOCH (1872) the *Nummulites* - *Discocyclus* limestone can be divided into three parts, which are the following from the bottom to the top:

- *Nummulites* limestone
- sandy limestone
- *Discocyclus* limestone.

Thick strata of the light brown *Nummulites* limestone are cropping out at 7 m of thickness. The uppermost one metre is oolitic with fine lime mud matrix, carbonate content is 70-100 % (MONOSTORI, 1967).

This rock is rich in coralline algae, *Nummulites* and *Miliolina* species.

The dominant bivalve species are *Chlamys*, *Plicatula bovensis* DE GREGORIO and *Lentipecten corneus* (SOWERBY). *Spondylus* and *Ostrea* species are abundant as well. There are many Echinoidea in the upper part of *Nummulites* limestone. The most frequent species are *Echinanthus scutella* LAMARCK, *Echinolampas subsimilis* D'ARCHIAC and *Sismondia rosacea* LESKE.

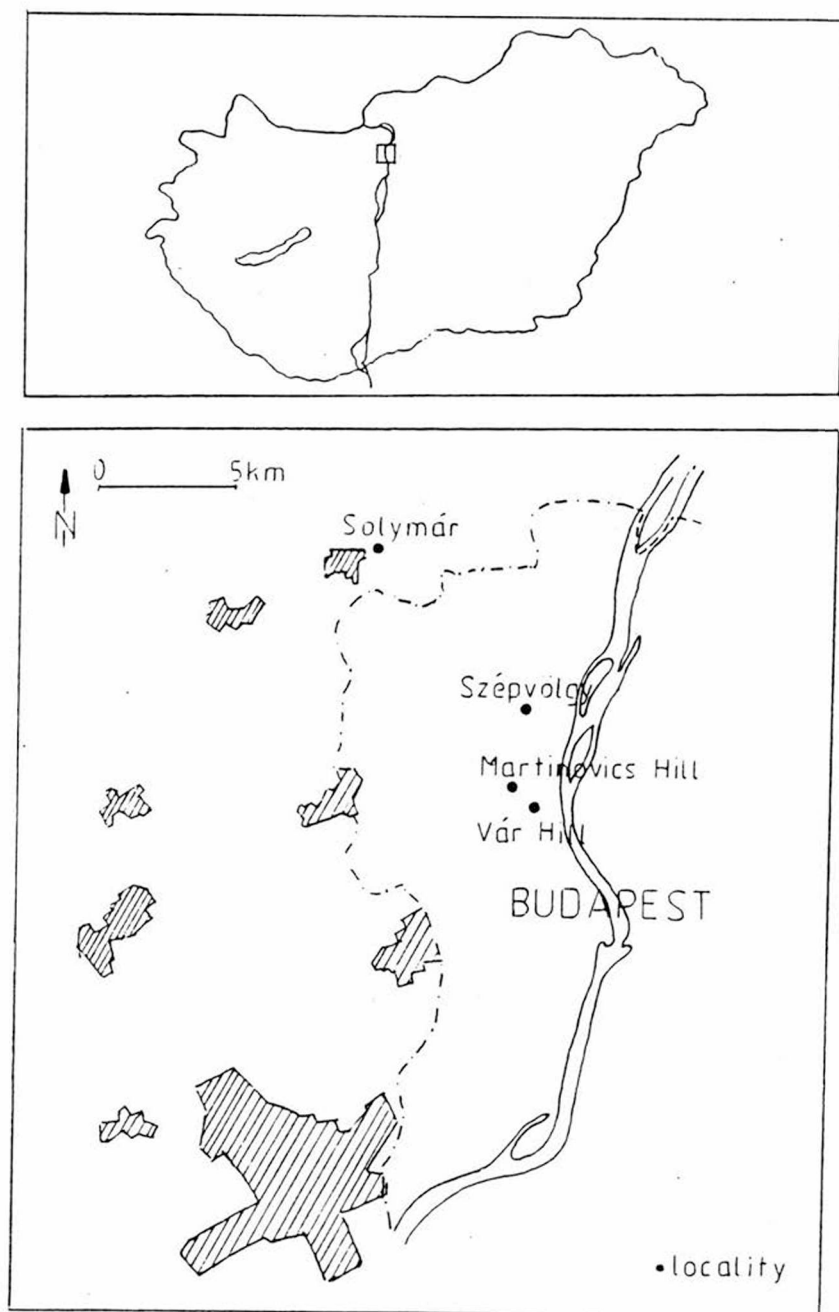


Fig. 1. Upper Eocene mollusc localities in Buda Mts.

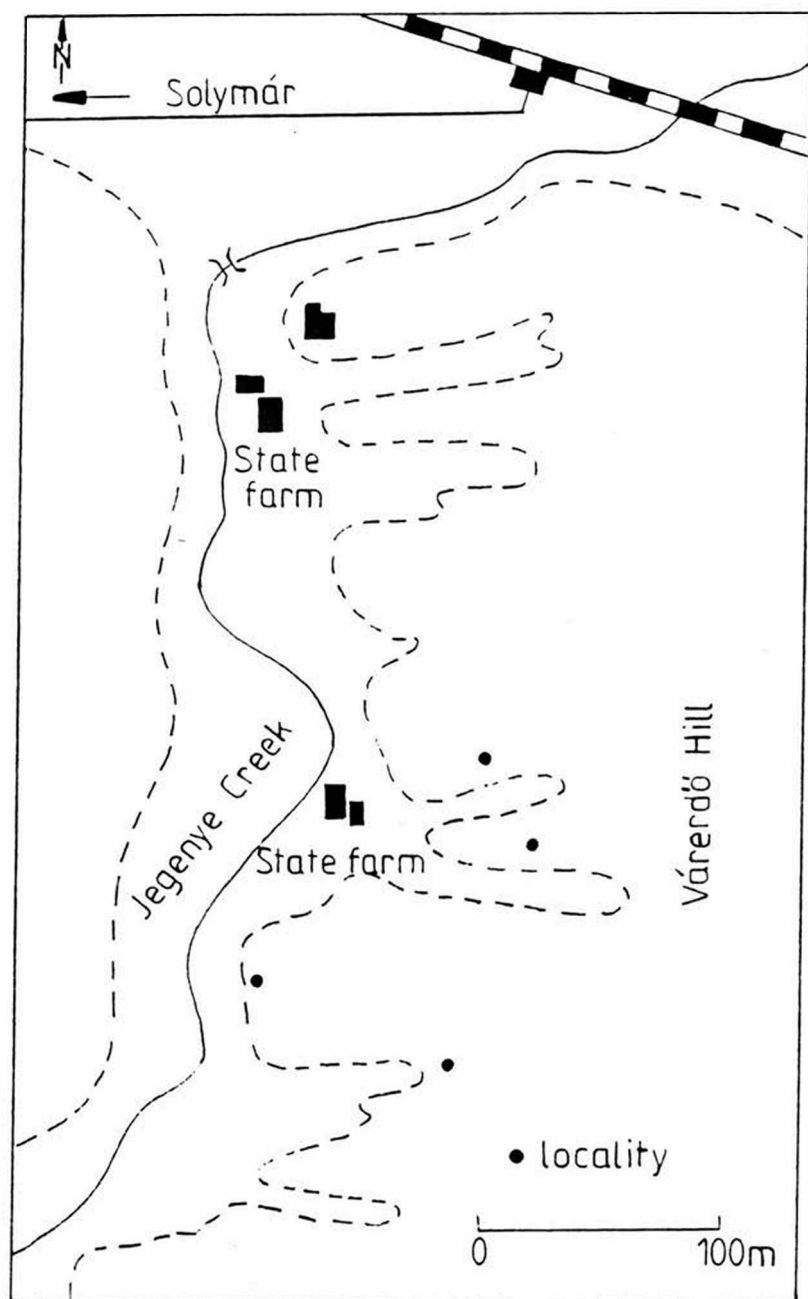


Fig. 2. Localities at Solymár, Várerdő Hill

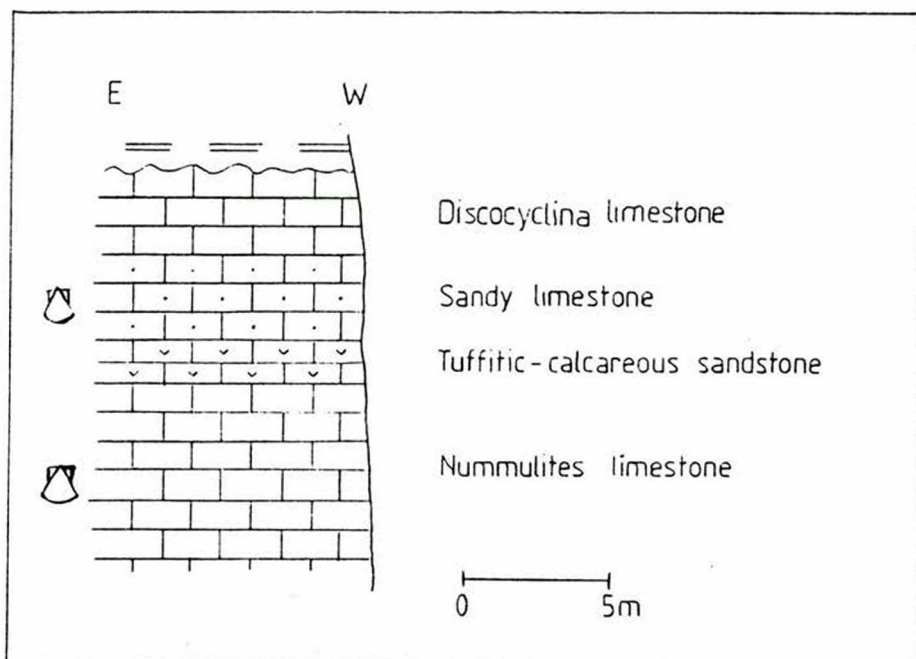


Fig. 3. Upper Eocene sequence at Solymár

The *Nummulites* limestone is gradually changed into the lilac - red, 6 m thick Sandy limestone, in which lime grains become dominant in red, argillaceous, calcareous muddy matrix. The carbonate content is 70 % (MONOSTORI, 1967).

Coralline algae are less important, while *Discocyclus* is more abundant.

More *Chlamys biarritzensis* (D'ARCHIAC) and *Plicatula bovensis* and more *Echinanthus scutella*, *Sismondia rosacea* and *Scutella tenera* LAUBE can be found.

The sandy limestone is followed by a 2-3 m thick, tabular *Discocyclus* limestone free of macrofossils.

2. Szépvölgy

The highest abundance of bivalves is in the Bryozoa Marl of Szépvölgy.

2.1 Western quarry of Mátyás Hill

It is in front of Pálvölgy Cave (Fig. 4).

KÁZMÉR (1982) distinguished the following:

- *Orthophragmina* limestone (0-10 m)
- Calcareous *Actinocyclus* marl (10-13 m)
- Grey *Actinocyclus* marl (13-15 m)
- Bryozoan Marl (15-30 m) with an intercalation of quartzose sandstone (0.2 m)
- Buda Marl.

2.2 Eastern quarry of Mátyás Hill

This quarry is situated on the southern slope of the hill, near to Mátyáshegy and Kolostor streets.

MONOSTORI (1965) described the following units upward:

- limestone with coralline algae
- *Discocyclus* limestone
- Bryozoan Marl.

2.3 Outcrop at Erdeilak Restaurant

The Bryozoan Marl crops out 100 m westward from the restaurant.

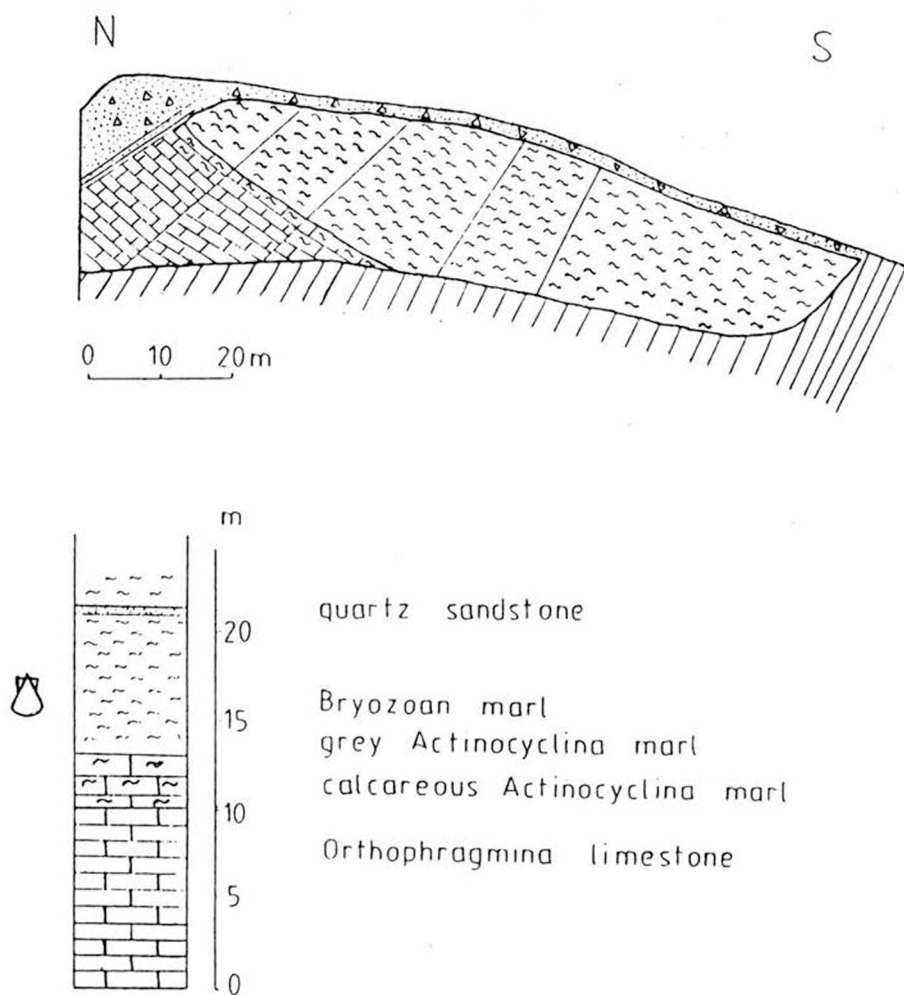


Fig. 4. Mátyás Hill, western quarry. Sequence after KÁZMÉR, 1982.

In this three last localities the Bryozoa Marl contained the bivalve fauna. The light grey, silty marl is rich in fossils. The dominant foraminifer genera are *Actinocyclus* and *Discocyclus*. Bryozoa are rock-forming. The 75 % of the bivalve fauna is *Chlamys biarrizensis*. From the rich Echinoidea assemblage *Schizaster lorioli* PÁVAY and *Opissaster szecsenyii* (PÁVAY) are the most frequent species.

Considerable part of the specimens were gathered from presently unknown outcrops of Martinovcs Hill, Vár Hill and Némervölgy road.

3. Martinovics Hill

Ninety specimens were collected from the *Nummulites* - *Discocyclus* limestone. Although this assemblage is smaller in abundance, it is greater in number of species. The 90 specimens belong to 10 species.

4., 5. Vár Hill and Némervölgy road

Twelve individuals from Vár Hill and sixteen from Némervölgy road were collected in the Buda Marl from house foundations.

General character of the bivalve fauna

The 370 specimens of this bivalve fauna belong to 21 species of 15 genera. Moulds and embedded shells made the determination more difficult, therefore the internal morphological features could not be studied. Many shells were etched so the external ornaments were missing. Due to the bad preservation 130 specimens were ranged only into genera.

Our determinations are in good agreement with the system of Treatise (MOORE, 1969). Table 1. contains the complete fauna list.

Table 1. Complete list of the studied fauna

Phylum Mollusca

Classis Bivalvia LINNÉ, 1758

Subclassis Pteriamorphia BEURLÉN, 1944

Ordo Arcoida STOLICZKA, 1871

Superfamilia Limopsacea DALL, 1895

Familia Glycymerididae NEWTON, 1922

Subfamilia Glycymeridinae NEWTON, 1922

Genus *Glycymeris* DA COSTA, 1778

Glycymeris sp.

Ordo Mytiloida FÉRUSAC, 1822

Superfamilia Mytilacea RAFINESQUE, 1815

Familia Mytilidae RAFINESQUE, 1815

Subfamilia Lithophaginae ADAMS and ADAMS, 1857

Genus *Lithophaga* RÖDING, 1798

Lithophaga zignoi (OPPENHEIM), 1900-1901

Subfamilia Modiolinae KEEN, 1958

Genus *Modiolus* LAMARCK, 1799

Modiola modioloides (BELLARDI), 1852

Modiola cf. *subcarinata* LAMARCK, 1806

Superfamilia Pinnacea LEACH, 1819

Familia Pinnidae LEACH, 1819

Genus *Pinna* LINNÉ, 1758

Pinna cf. *margaritacea* LAMARCK,

Ordo Pterioda NEWELL, 1965

Subordo Pteriina NEWELL, 1965

Superfamilia Pectinacea RAFINESQUE, 1815

Familia Pectinidae RAFINESQUE, 1815

Genus *Lentipecten* MARWICK, 1928

Lentipecten corneus (SOWERBY), 1821

Genus *Propeamussium* DE GREGORIO, 1884

Propeamussium semiradiatus (MAYER), 1861

Parvamussium fallax (KOROBKOV), 1939

Genus *Chlamys* RÖDING, 1798

Chlamys biarritzensis (D'ARCHIAC), 1847

Chlamys aff. *multicarinata* (DESHAYES), 1824

Chlamys subdiscors (D'ARCHIAC), 1847

Familia Spondylidae GRAY, 1826

Genus *Spondylus* LINNÉ, 1758

Spondylus bifrons MÜNSTER, 1840

Spondylus buchi PHILIPPI, 1847

Spondylus cf. *cisalpinus* BRONGNIART, 1823

Spondylus radula LAMARCK, 1806

Familia Plicatulidae WATSON, 1930

Genus *Plicatula* LAMARCK, 1801

Plicatula bovensis DE GREGORIO, 1894

Subordo Ostreina FÉRUSAC, 1822

Superfamilia Ostreacea RAFINESQUE, 1815

Familia Grypheidea VYALOV, 1936

Subfamilia Pycnodontinae STENZEL, 1959

Genus *Pycnodonta fischer* DE WALDHEIM, 1835

Pycnodonta brongniarti (BRONN), 1831

Familia Ostreidae LAMARCK, 1818

Subfamilia Ostreinae RAFINESQUE, 1815

Genus *Cubitostrea* SACCO, 1897

Cubitostrea cf. *plicata* SOLANDER,

Subfamilia Lophinae VYALOV, 1936

Genus *Lopha* BOLTON in RÖDING, 1798

Lopha martinsi (D'ARCHIAC), 1848

Classis Heterodonta NEUMAYR, 1884

Ordo Veneroida ADAMS and ADAMS, 1856

Superfamilia Lucinacea FLEMING, 1828

Familia Lucinidea FLEMING, 1828

Subfamilia Lucininae FLEMING, 1828

Genus *Lucina* BRUGUIÈRE, 1797

Lucina sp.

- Superfamilia Carditacea FLEMING, 1820
- Familia Carditidae FLEMING, 1828
- Subfamilia Carditesianae CHAVAN, 1966
- Genus *Cardites* LINK, 1807

Cardites sp.

- Superfamilia Crassatellacea FÉRUSAC, 1822
- Familia Crassatellidae FÉRUSAC, 1822
- Subfamilia Crassatellinae FÉRUSAC, 1822
- Genus *Crassatella* LAMARCK, 1799

Crassatella curata DESHAYES,
Crassatella subtumida BELLARDI, 1852

- Subclass Anomalodesmata DALL, 1889
- Ordo Pholadomyoida NEWELL, 1965
- Superfamilia Pholadomyacea GRAY, 1847
- Familia Pholadomyidae GRAY, 1847
- Genus *Pholadomya* SOWERBY, 1823

Pholadomya loczyi TAEGER, 1909

Biostratigraphy

For drawing biostratigraphical conclusions in the Late Eocene, plankton foraminifer and nannoplankton assemblages are more suitable than bivalves. The nannoplankton studies were carried out by BÁLDI-BEKE (1970).

Buda Marl containing *Variamussium fallax* is comparable with *Variamussium fallax* zone of Late Eocene described by Soviet stratigraphers.

The presence of this species let more precise correspondence between Paleogene of Buda and Crimea, Caucasus and Rhodope regions (BÁLDI, 1983).

The great bulk of bivalves from Buda Hills indicate Late Eocene, but some of them existed in Middle Eocene or in Oligocene as well. Species confining only to the Middle Eocene or the Oligocene were not identified.

Ranges of identified species are shown in Table 2.

Table 2. Ranges of identified species

	E ₂	E ₃	Ol ₁
<i>Chlamys biarritzensis</i> (D'ARCH.)		x	x
<i>Chlamys subdiscors</i> (D'ARCH.)	x	x	x
<i>Crassatella curata</i> DESH.	x	x	
<i>Crassatella subtumida</i> BELL.	x	x	
<i>Cubitostrea</i> cf. <i>plicata</i> SOL.	x	x	
<i>Lentipecten corneus</i> (SOW.)	x	x	x
<i>Lithophaga zignoi</i> (OPP.)		x	
<i>Lopha martinsi</i> (D'ARCH.)	x	x	x
<i>Modiola modioloides</i> (BELL.)		x	
<i>Modiola</i> cf. <i>subcarinata</i> LAM.	x	x	
<i>Parvamussium fallax</i> (KOR.)	x	x	x
<i>Pholadomya loczyi</i> TAEG.	x	x	
<i>Pinna</i> cf. <i>margaritacea</i> LAM.	x	x	
<i>Plicatula bovensis</i> DE GREG.		x	
<i>Propeamussium semiradiatus</i> (MAY.)		x	x
<i>Pycnodonta brongniartii</i> (BRONN)	x	x	x
<i>Spondylus bifrons</i> MÜNST.	x	x	x
<i>Spondylus buchi</i> PHIL.	x	x	x
<i>Spondylus</i> cf. <i>cisalpinus</i> BRONGN.	x	x	x
<i>Spondylus radula</i> LAM.	x	x	

Palaeobiogeography

Twenty one species of the studied fauna were suitable for palaeobiogeographical reconstruction. There were only few comparable assemblages elaborated from different points of view or show different ecological or stratigraphical features, so the reconstruction was difficult to do. Statistical evaluation could not be carried out, as well.

Monographs with comparable faunal assemblages are only about the Transylvanian (MÉSZÁROS, 1957), Priabonian (PICCOLI et MOCELLIN, 1962) and Biarritzian (BOUSSAC, 1911) Late Eocene molluscs.

Taking into consideration the number of species in common of these four localities (Table 3), preliminary conclusions could be done.

This study, corresponding to KECSKEMÉTI (1980), proves that our bivalve assemblage belongs to the Mediterranean province.

Table 3. Number of species in common

	Number of species	Number of species in common
Buda Hills	21	—
Priabona	62	12
Transylvania	65	11
Biarritz	34	5

After Piccoli et Mocellin, 1962 ; Mészáros 1957; Boussac, 1911.

Palaeoecology

Study of fossils is useful in reconstruction of palaeoenvironments as well as in drawing conclusions about the conditions of sedimentation. In addition we can get information about the habitat of animals from the characters of sediments (AGER, 1963). In this paper both methods were used.

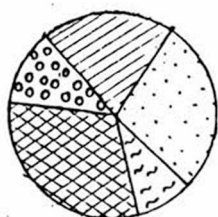
Studying the recent environmental circumstances: water depth, quality of sediments, temperature, salinity and hydrodynamics, we can draw conclusions about ancient life. It made the actualistic evaluations easy that 13 from the 15 identified genera are still living in the seas.

Our knowledge of the recent life is due to the excellent description of DAVITASHVILI and MERKLIN (1966). Not only the environmental demands of recent forms, but also the sedimentological and palaeontological data were taken into consideration.

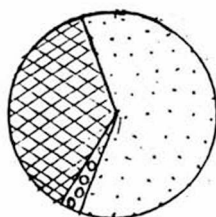
Fig. 5. shows the distribution of the characteristic bivalve genera of the main rock types. *Chlamys* is the dominant genus in *Nummulites* and sandy limestones of Solymár, in *Nummulites* - *Discocyclus* limestone and Bryozoan Marl. In the Buda Marl the most abundant genus is *Propeamussium*.

The palaeoecological conclusions were taken concerning the different abundance and number of species in the five rock types (Table 4).

Nummulites limestone

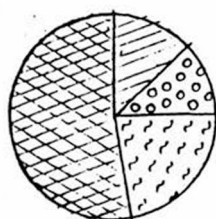


Sandy limestone

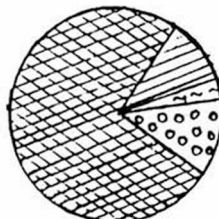


Nummulites

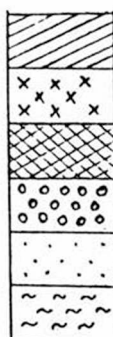
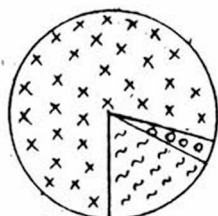
-Discocyclus limestone



Bryozoan marl



Buda Marl



Lentipeecten

Propeamussium

Chlamys

Spondylus

Plicatula

Other

Fig. 5. Distribution of 6 bivalve genera in 5 types of rock.

Table 4. Occurrences of species in the five rock types

	A	B	C	D	E
<i>Cardita</i> sp.			x		
<i>Chlamys biarritzensis</i> (D'ARCH.)	x	x	x	x	
<i>Chlamys</i> aff. <i>multicarinata</i> (DESH.)	x				
<i>Chlamys subdiscors</i> (D'ARCH.)			x		
<i>Chlamys</i> sp.	x	x	x	x	
<i>Crassatella curate</i> DESH.			x		
<i>Crassatella subtumida</i> BELL.			x		
<i>Cubitostrea</i> cf. <i>plicata</i> SOL.			x		
<i>Glycymeris</i> sp.					x
<i>Lentipecten corneus</i> (SOW.)	x			x	x
<i>Lithophaga zignoi</i> (OPP.)			x		
<i>Lopha martinsi</i> (D'ARCH.)			x		
<i>Lucina</i> sp.			x		
<i>Modiola modioloides</i> (BELL.)				x	
<i>Modiola</i> cf. <i>subcarinata</i> LAM.			x	x	
<i>Parvamussium fallax</i> (KOR.)					x
<i>Pholadomya loczyi</i> TAEG.					x
<i>Pinna</i> cf. <i>margaritacea</i> LAM.				x	
<i>Plicatula bovensis</i> DE GREG.	x	x			
<i>Propeamussium semiradiatus</i> (MAY.)					x
<i>Pycnodonta brongniarti</i> (BRONN)	x			x	
<i>Spondylus bifrons</i> MÜNST.				x	
<i>Spondylus buchi</i> PHIL.			x		
<i>Spondylus</i> cf. <i>cisalpinus</i> BRONGN.					x
<i>Spondylus radula</i> LAM.	x	x	x	x	

A *Nummulites* limestone

B Sandy marl

C *Nummulites-Discocyclus* limestone

D Bryozoan marl

E Buda Marl

1. *Nummulites limestone*

Besides the dominant *Chlamys*, *Picatula bovensis* and *Lentipecten corneus*, *Spondylus* and *Ostrea* are also very frequent. Recent *Chlamys* species live in warm, shallow water attached by byssal threads to the sandy bottom or swim free. *Lentipecten* with smooth, flat shells is supposed to have lived in crevices in still water. Thick shells of *Ostrea* and spines of *Spondylus* indicate currents in the shallow water.

2. *Sandy limestone*

While *Picatula bovensis* and *Chlamys* are very frequent, *Lentipecten corneus* is absolutely missing in this rock type. Although the bivalve fauna of *Nummulites* limestone and sandy limestone differ from each other, significant differences in the environmental conditions cannot be recognized.

3. *Nummulites-Discocyclina limestone*

Although the fauna of this rock type is the richest, more than the half of the specimens are *Chlamys*.

The morphology of the depositional surface must have been variable. *Lucina* lived in the sandy mud of shoals. *Lithophaga* bored into calcareous formations with the aid of acid secretion. The mollusc was held in position by byssal threads attached to the side of the burrow. The modern *Cardita* is abundant on the sandy bottom, above 50 m, in strong currents. Cemented forms also occur (eg.: *Spondylus*). According to recent examples this fauna lived on the sandy or muddy bottom, deeper than the wave base, but in maximum depth of 60 m, in warm water of normal salinity.

4. *Bryozoan marl*

Chlamys biarritzensis gives the 75 % of this assemblage. The remaining specimens are mostly *Spondylus* or *Lentipecten corneus*.

These genera refer to a wide range of water depths. Modern *Chlamys* species points to 1-90 m of depth. *Pinna* occurs from 30 to 300 m, in the soft, fine sand. *Pinna* is partly buried in sea bottom with pointed anterior end lowermost, attached by byssus to rocks or other objects in the sediment.

Spondylus appear even beneath 200 m, while *Modiola* lives between littoral and pelagic zones.

Interpreting this palaeocommunity we can suppose that Bryozoan Marl was deposited in approximately 100 m of warm, normal salinity water.

5. Buda Marl

The fauna gathered from Buda Marl indicates the deepest environment.

Amussium, preferring deep, still water substitutes *Chlamys*, which dominates the previous rocktype. The attending *Glycymeris* and *Pholadomya* refer to deep water, too. In addition the lack of spines and ribs on shells points to deep, still water. Buda Marl may have been deposited in few hundred metres of depth.

In this sequence the bivalve assemblage indicate the deepening sea. The bottom sediments become finer from sand to mud.

Acknowledgements

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REFERENCES

- AGER, D. V. (1963): Principles of Paleogeology - New York, McGraw Hill, 371. p.
- BALÁZS E. - BÁLDI T. - DUDICH E. - GIDAI L. - KÖRPÁS L. - RADÓCZ GY. - SZENTGYÖRGYI K. - ZELENKA T. (1981): A magyarországi eocén-oligocén határ képződményeinek szerkezeti-faciális vázlata (Structural and facies outline of Eocene/Oligocene boundary formations in Hungary (in Hungarian) - Földtani Közöny 111/1. pp. 145-156. 8 fig.
- BÁLDI T. (1986): Mid-Tertiary stratigraphy and paleogeographic evolution of Hungary (Magyarországi oligocén és alsómiocén formációk) - Akadémiai Kiadó, 293 p. Budapest
- BÁLDI - BEKE M. (1970): The Nannoplankton of the Bryozoan and Buda Marls (Palaeogene of Budapest, Hungary) (Hung. with English abstract) - Őslénytani viták 16. pp. 31-50., 5 fig.
- BOKOR GY. (1939): A Budai-hegység nyugati peremének földtani viszonyai (The Geology of the Western Border of the Mountains of Buda) (in Hungarian) - Földtani Közöny 69/10-12. pp. 219-233., 4 fig.

- BOUSSAC, J. (1911): Études stratigraphiques et paléontologiques sur le Nummulitique de Biarritz - Ann. Hébert tom. V. pp. 1-94., 24 Pl.
- DAVITASHVILI, L. S., MERKLIN, R. L. (eds.) (1966): Spravochnik po ekologiji morskikh dvustvorok - Nauka, Moskva, 352 p.
- DUDICH, E. jr. (1959): Paläogeographische und paläobiologische Verhältnisse der Budapester Umgebung in Obereozän und Unteroligozän - Ann. Univ. Sci. Budapest, Sect. Geol. 2. pp. 53-87. 4 fig.
- HOFMANN K. (1873): Beiträge zur Kenntniss der Fauna des Haupt-Dolomites und der älteren Tertiär-Gebilde des Ofen-Kovácsier Gebirges - Földt. Int. Évk. II. 26 p. 5 Tav. Budapest
- KÁZMÉR M. (1982): A budai felsőeocén mészkő mikrofácies vizsgálata (Unpublished doctoral thesis) - (Eötvös University, Department of Geology, Budapest, I-II. 144. p.)
- KECSKEMÉTI T. (1980): A Bakony hegységi Nummulites fauna paleobiográfiája (Aperçu paléobiogéographique sur la faune de Nummulites du Bakony) (in Hung. with French abstract) - Földt. Közl. 110/3-4. pp. 432-445.
- KOCH A. (1872): A Solymár melletti Várerdőhegy földtani viszonyai (in Hungarian) - Földt. Közl. I., pp. 93-95.
- KOPEK G. - KECSKEMÉTI T. - DUDICH E. (1966). A Dunántúli-Középhegység eocénjének rétegtani kérdései (Stratigraphische probleme des Eozäns im Transdanubischen Mittelgebirge) (in Hung. with German abstract) - Magyar Állami Földtani Intézet Évi Jelentése 1964-évi, pp. 246-264.
- LŐRENTHEY I. (1897): Paleontológiai tanulmányok a harmadkorú rákok köréből (in Hungarian) - Math. Term. Tud. Közl. 18/2, pp. 1-16.
- LŐWY B. (1928): A budai Kis-Svábhegy földtani viszonyai (in Hungarian) (Doctoral thesis) 30. p. Budapest
- MÉSZÁROS N. (1957): Fauna de molute a depozitelor paleogene din Nord-Vestul Transilvaniei - Acad. Rep. Pop. rom p. 174, 20 tav. Cluj
- MONOSTORI M. (1965): Paläoökologische und Faziesuntersuchungen an den Obereozän-Schichten in der Umgebung von Budapest - Ann. Univ. Sci. Budapest Sect. Geol. 8. pp. 139-149. 1 fig. 3 Taf
- MONOSTORI M. (1967): Paläogene Faziesuntersuchungen am Várerdő-Berg bei Solymár - Ann. Univ. Sci. Budapest Sect. Geol. 10. pp. 161-176. 5 fig. 1 Taf.
- MOORE, C. R. (1969): Treatise on Invertebrate Paleontology Part N., Mollusca 6. Bivalvia - Geol. Soc. America, Univ. Kansas Press
- PICCOLI-MOCELLIN (1962): Studi sulla macrofauna priaboniana di Priabona (Prealpi Venete) - Padova Soc. Coop. Tip. 120 p. 5 Tav.
- WEIN Gy. (1977). A Budai-hegység tektonikája - MÁFI, 76 p. (in Hungarian).